UNITED STATES DEPARTMENT OF THE INTERIOR GEOLOGICAL SURVEY

REGIONAL GEOHYDROLOGY OF THE NORTHERN LOUISIANA SALT-DOME BASIN,

PART I, CONCEPTUAL MODEL AND DATA NEEDS

By G. N. Ryals

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JAMES G. WATT, Secretary

GEOLOGICAL SURVEY

Dallas L. Peck, Director

For additional information write to:

District Chief U.S. Geological Survey, WRD P.O. Box 66492 Baton Rouge, LA 70896 (504) 389-0281 For sale by:

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FACTORS FOR CONVERTING INCH-POUND UNITS TO INTERNATIONAL SYSTEM (SI) OF METRIC UNITS

Multiply	<u>By</u>	To obtain
foot (ft)	0.3048	meter (m)
foot per day (ft/d)	0.3048	meter per day (m/d)
foot squared per day (ft^2/d)	0.9290	meter squared per day (m^2/d)
gallon per minute (gal/min)	0.06309	liter per second (L/s)
	3.785×10^{-3}	cubic meter per minute (m ³ /min)
inch (in.)	25.40	millimeter (mm)
mile (mi)	1.609	kilometer (km)
square mile (mi ²)	2.590	square kilometer (km ²)

To convert temperature in degree Celsius (°C) to degree Fahrenheit (°F), multiply by 9/5 and add 32.

REGIONAL GEOHYDROLOGY OF THE NORTHERN LOUISIANA SALT-DOME BASIN, PART I, CONCEPTUAL MODEL AND DATA NEEDS

By G. N. Ryals

ABSTRACT

As part of the National Waste Terminal Storage Program, the U.S. Geological Survey is conducting a regional study of the geohydrology of the northern Louisiana salt-dome basin and developing a regional multilayered ground-water flow model to determine regional flow paths.

In the salt-dome basin the Tokio Formation and Brownstown Marl (Austin aquifer in this report), and Nacatoch Sand of Late Cretaceous age and the Wilcox Group, Carrizo Sand, Sparta Sand, and Cockfield Formation of Tertiary age contain regional aquifers within the maximum potential repository depth of 3,000 feet. The Cretaceous units contain saltwater throughout the basin. The Tertiary units contain freshwater to varying distances downdip from outcrop areas in the basin. Natural flow directions and rates of movement of ground water have been changed in the salt-dome basin by the withdrawal of freshwater and by the injection of wastes (principally oil-field brines) into saline aquifers. Except for the Sparta aquifer, ground-water flow directions are not well known because of a lack of potentiometric data.

A regional test-drilling program, to collect the data needed to document concepts of the flow system and to quantify inputs to the planned ground-water flow model, has been proposed. The Sparta aquifer is being modeled first because data are available for the unit. As regional test drilling provides data on other units, those units will be added to the model developed for the Sparta aquifer.

INTRODUCTION

The Department of Energy (DOE), formerly the Energy Research and Development Administration, in 1976 began an expanded waste-management program for both defense and commercially produced radioactive waste. The National Waste Terminal Storage (NWTS) program is an effort by DOE to locate and develop sites in various parts of the country for disposal or storage of commercially produced radionuclides in deeply buried geologic formations. The Office of Nuclear Waste Isolation, Battelle Memorial Institute, at Columbus, Ohio, administers the NWTS program for DOE. As

part of the program, salt domes in the Gulf Coast Region of Texas, Mississippi, and Louisiana (the northern Louisiana salt-dome basin) are being considered for their suitability as repositories. Law Engineering Testing Co. (LETCO) served as the Geologic Project Manager coordinating study activities and data acquisition for the Gulf Coast Region, 1977-81. In Louisiana, the U.S. Geological Survey's participation in the NWTS Program, in cooperation with DOE, has been to describe the regional geohydrology of the northern Louisiana salt-dome basin and to develop a regional multilayered ground-water flow model. This report presents a brief conceptual model of the ground-water flow system of the northern Louisiana salt-dome basin, a discussion of the modeling effort, and outlines data needs to accomplish the modeling.

Location

The northern Louisiana salt-dome basin has an area of about 3,000 mi² and includes all or parts of 11 parishes in north-central and northwestern Louisiana. The area of interest for the ground-water flow model shown in figure 1 is considerably larger than the northern Louisiana salt-dome basin, as most of the aquifers have regional extent, and ground-water flow within the basin is part of regional patterns. Model boundaries will extend beyond the principal area of interest to incorporate aquifer boundary conditions.

Repository Site Selection

The Office of Nuclear Waste Isolation, (ONWI) in the overall program to select a repository site, is using a four-phase approach in which successively more detailed studies are conducted on successively smaller study areas. The first phase consists of regional broad-based evaluations of multistate regions to identify areas that may be suitable for repositories. In the second phase, area studies of about 1,000 mi² lead to the selection of locations for further study. The location phase consisting of approximately 30 mi² study areas will lead to the identification of a site for an exploratory shaft of a potential repository site or sites in a region. Finally, these potential sites will be studied in detail (10 mi² study areas), and one or more sites will be chosen for a test facility leading ultimately to a repository.

In the Gulf Coast Region studies, the four characterization phases are called Regional, Area, Location, and Detailed Site Characterization. In the Regional Characterization studies, 8 of 263 known salt domes were chosen for study in the Area Characterization phase (Law Engineering Testing Company, 1978). Rayburns and Vacherie salt domes in Louisiana (fig. 2) were two of the eight domes chosen for study in the Area phase. The Area Characterization studies, completed in 1980 (Law Engineering Testing Company, 1980a, b, c, d), provided the data for ONWI to select four of the eight salt domes (Office of Nuclear Waste Isolation, 1980) for further study in the Location Characterization phase. Vacherie dome is one of the four domes selected. Currently, plans for Location and Detailed Site Characterization studies are being developed.



Figure 1.--Location of northern Louisiana salt-dome basin.

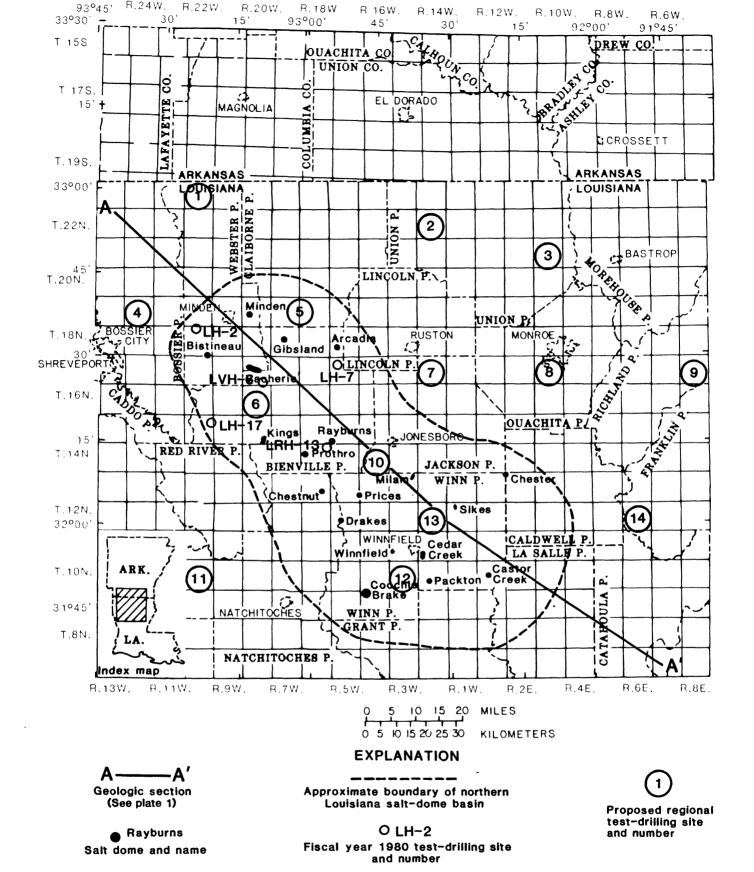


Figure 2.--Location of salt domes, fiscal year 1980 test-drilling sites, and proposed regional test-drilling sites.

The four-phase approach for selecting repository sites seems to be advantageous for geologic formations of large areal extent. Granite, tuff, basalt, and bedded-salt are rock types of large areal extent currently being considered as hosts for a repository. For these rock types, the focus is from regional studies to the most suitable site for a potential repository in the formation as the four-phase program proceeds. As the study areas decrease in size, the regional-flow system is studied in less detail and the local flow system in more detail. However, in formations having large areal extent, the local ground-water flow regime probably would be an integral part of the regional flow system, whereas the salt domes in the Gulf Coast probably are not.

Salt domes are the only geologic formations of small areal extent that are being considered as repositories. During dome development, salt domes have pierced thousands of feet of Cretaceous and Tertiary sediments that contain thick aquifers. Many of these are major aquifers that extend into several states. In a regional setting, the salt domes appear as local anomalies. As a result of the approach of studying successively smaller areas, more detailed attention is given to these anomalous areas for comparing one to another. Because of the complexity of domal areas, localized detailed studies are needed. The regional flow system in the vicinity of the salt domes and related structures has been modified. Therefore, regional as well as local flow systems must be known in As pointed out by Hosman (1978), both regional and local geohydrologic conditions will control movement should a contaminant escape from a salt dome.

Participation by the Geological Survey

In this investigation, the objective of the U.S. Geological Survey in Louisiana is to study the regional geohydrology. The Survey has not participated in the selection of domes for further study, nor does the Survey plan to participate in selecting a specific dome for a potential However, the Survey has cooperated with ONWI and LETCO during the Regional and Area Characterization studies by providing data and preparing specific reports needed for the site selection process. During the Regional Characterization phase, the general geohydrology of the northern Louisiana salt-dome basin was described by Hosman (1978) using existing data. As part of the Area Characterization phase, completed in 1980, a total of 16 test wells were completed at 5 sites in Louisiana (fig. 2). The test-drilling program was planned and managed by LETCO; however, the Survey participated in the selection of sidewall-core depths and screen intervals, sampled test wells for water-quality determinations, and analyzed aquifer tests. Maps of the base of fresh ground water and the potentiometric surface of the Sparta and Wilcox-Carrizo aquifers were prepared (Ryals, 1980a, b, c). In addition, hydrologic data from the vicinity of the Vacherie and Rayburns salt domes were compiled (Ryals and Hosman, 1980).

CONCEPTUAL MODEL OF AQUIFER SYSTEM

General Geohydrologic Framework

The main structural feature in north-central Louisiana is the north Louisiana syncline (fig. 1). The syncline encompasses two structural basins (evident on structure maps of Cretaceous units) that are referred to as the northern Louisiana salt-dome basin (fig. 2) because of the presence of 19 known salt domes. The basin is bounded on the east by the Monroe uplift and on the west by the Sabine uplift. The southern part of the basin is bounded by the Angelina-Caldwell flexure, which is at the northern edge of Miocene deposits. To the north the basin becomes less distinct.

A repository in a salt dome probably will be within a depth range of 1,000 to 3,000 ft below land surface (Office of Nuclear Waste Isolation, 1981). Geologic units above 3,000 ft in the northern Louisiana salt-dome basin are of Late Cretaceous, Tertiary, and Quaternary age (table 1). The Upper Cretaceous units consist of marine chalk, marl, limestone, The Tertiary units consist of alternating marine clays clay, and sand. and nonmarine sands. The Cretaceous and Tertiary units generally dip and thicken to the southeast (pl. 1). The Quaternary deposits consist of clay, sand, and gravel and are nearly flat lying. They were deposited on an irregular, eroded Tertiary surface. Table 1 gives a brief description of the lithology and an indication of the water-bearing characteristics of the geologic units. Detailed discussions of regional geology hydrology can be found in Boswell and others (1965), Cushing and others (1964 and 1970), Hosman and others (1968), Hosman (1978), and Payne (1968, 1970, 1972, and 1975).

In 1980 as part of the Area Characterization studies, a total of 16 test wells were completed at five test-drilling sites (fig. 2) and the data collection needed for describing the regional as well as near-dome hydrology began. Sites LVH-6, LH-7, and LRH-13 are near salt domes. Data collected from these sites are probably more indicative of near-dome hydrology than of regional hydrology. However, data from sites LH-2 and LH-17 are indicative of regional conditions. Table 2 gives the results of aquifer tests and other pertinent information for the 16 test wells, and chemical analyses of water samples from various aquifers are given in tables 4, 5, and 6. Detailed presentations of the data collected from the Area Characterization test-drilling activities were published as well-completion reports by Law Engineering Testing Company, Marietta, Ga. (1981a, b, c, d, e).

In the salt-dome basin between a depth of 3,000 ft and land surface, six geologic units contain regional aquifers. From oldest (deepest) to youngest, the aquifers are in the Tokio Formation and Brownstown Marl (Austin aquifer in this report), and Nacatoch Sand of Late Cretaceous age, and the Wilcox Group, Carrizo Sand, Sparta Sand, and Cockfield Formation (southeastern part of basin) of Tertiary age. The Wilcox is hydraulically interconnected with the overlying Carrizo; therefore, the

Carrizo and the Wilcox are treated as one hydrologic unit, the Wilcox-Carrizo aquifer. The aquifers are separated by confining layers that retard water movement. Water in each unit is confined under artesian pressure, except in the outcrop where water-table conditions prevail.

In general, under natural-flow conditions, water moves from points of recharge in the outcrop areas downdip to discharge areas in the subsurface. In discharge areas water moves upward into the overlying units because the head increases with depth.

Water enters the Cretaceous aquifers in outcrop areas in southern Arkansas and eastern Texas. Discharge areas for the Cretaceous aquifers have not been well established. Because the sand facies of the Nacatoch Sand (the Nacatoch Sand aquifer) (fig. 3) occurs only in the northwestern part of the project area, that area probably is the discharge area for the unit. The Tertiary aquifers crop out discontinuously in the saltdome basin and north of the basin. Water discharges from these units into the Mississippi River alluvial aquifer. In the saltdome basin, Quaternary deposits, which occur as alluvial fill and terrace remnants, are recharged locally by precipitation. In some areas, downward leakage from Quaternary deposits recharges Tertiary aquifers; but most of the discharge from the Quaternary deposits occurs by lateral movement to streams.

Saline water in the aquifers is flushed and displaced by freshwater moving downdip from recharge areas to replace water discharged vertically. Saline water has not been flushed from the Cretaceous sands in the study area but has been flushed to varying degrees in the Tertiary sands. Plate 2 shows the occurrence of freshwater in the study area and the extent of flushing in various units. In general, flushing has progressed from outcrop areas in the basin to about 20 mi downdip in the Wilcox-Carrizo aquifer and more than 100 mi downdip in the Sparta aquifer. Abrupt and large differences in the base of freshwater occur at the downdip limit of freshwater as the base changes from one aquifer to another.

Direction of Ground-Water Flow in the Major Aquifers

Man's activities in the salt-dome basin since the early 1900's have modified the natural flow system. The Wilcox-Carrizo and Sparta aquifers yield water for domestic, public-supply and industrial use; parts of the Wilcox-Carrizo, and the Nacatoch Sand and Austin aquifers are used for disposal of industrial wastes, principally oil-field brines. Some oil and gas is produced from the Austin. Recharge by injection wells has caused local potentiometric-high areas, but discharge (withdrawal) from wells in other areas has caused local and widespread cones of depression. The direction and rate of movement of ground water is controlled by the gradient of the potentiometric surface and by the geometry and hydrologic properties of the aquifers. The relatively simple gradients of the natural flow system have been modified to form a relatively complex flow system.

Table 1. --Generalized post-Lower Cretaceous stratigraphic column for northern Louisiana salt-dome basin and vicinity

Formation Formation Formation Formation Formation Formation Formation Fine lignitic sand and clay. Thickness 400 Claiborne Fine lignitic sand and carbon- Cockfield Fine lignitic sand and carbon- Lower part: Thickness 400 Clay, part ly sand and carbon- Linearbads: lignitic sand and carbon- Linearbads: lignitic sand and carbon- Linearbads: lignitic sand and carbon- Clay, part ly sand and slauconi- Clay, part ly sand and slauconi- Clay, part ly sand and slauconi- Lic. Thickness about 100 ft. Fine to medium sand with clay Rise to medium sand with clay Fine to medium sand with clay Fine to medium sand since so to 700 ft. Fine to medium sand since so to 150 ft. Fine to medium sand since so to 150 ft. Fine to medium sand since so to 150 ft. Fine to medium sand since so to 150 ft. Fine to medium sand since so to 150 ft. Linearbads: lignitic. Thick- Linuous. Thickness 0 to 150 ft. Fine to medium sand, clay, and Sol to 1,500 ft. Thickness ot 1,500 ft. Linearbaded sand, clay, and Sol to 1,500 ft. Thickness so to 1,500 ft. Thickness about 100 ft. Linearbaded sand, clay, and Sol to 1,500 ft. Thickness ot 1,500 ft.	Aqui- Hydrologic characteristics	Contains freshwater. Used locally for rural supplies and some public supplies. Yields range from a few gal/min for small domestic supplies to several thousand gal/min gal for large irrigation wells. Hydraulic conductivity ranges from 100 to 300 ft/d.	Contains freshwater and saltwater. Hydraulic conductivity of the ranges from about 25 ft/d to more than 100 ft/d.	Generally not water bearing. Local sands yield small		Contains freshwater and saltwater. Used mostly for small rural supplies. Hydraulic conductivity ranges from less than 15 ft/d to more than 40 ft/d.	Generally not water bearing. Local sands yield small quantities of water to wells.	Contains freshwater and saltwater. Principal aquifer of north-central Louisiana. Large withdrawals by domestic, municipal, and industrial wells. Hydraulic conductivity ranges from 30 ft/d to more than 100 ft/d.	Not water bearing.	Contains freshwater and saltwater. Penetrated mostly by small-yielding rural wells. Larger supplies developed	· · · · · · · · · · · · · · · · · · ·
valeo-cene Wilcox Wilcox Gartizo Sand Gartizo Gartizo	Description	alluvial valley graveliferous upward to sand, Thickness ft.				ne lignitic sand and caceous clay. Thicker lower part. Thickness 500 to 600 ft.	Clay, partly sandy and tic. Thickness about	Fine to medium sand with clay interbeds; lignitic. Thick-hanness 500 to 700 ft.	clay; some marl. 200 to 300 ft.		s s s s s s s s s s s s s s s s s s s
cene Eocene and Holocene and Holocene and Holocene Oligocene and Holocene On H			-ibnU	babivibnU	babivibnU	1	Moun- rain		River	рәрт	ΛŢΡUΩ
Eocene Holocene and Holocene and Holocene and Holocene and Holocene Cene	Group			N icksburg	ласкsоп			Claiborne)	 хо	MįĮo
Tertiary A S S S S S S S S S S S S S S S S S S		bns	Miocene	9nəsogi10				Еосепе			Paleo-
	Sys- tem	Quaternary						ertiary	T		

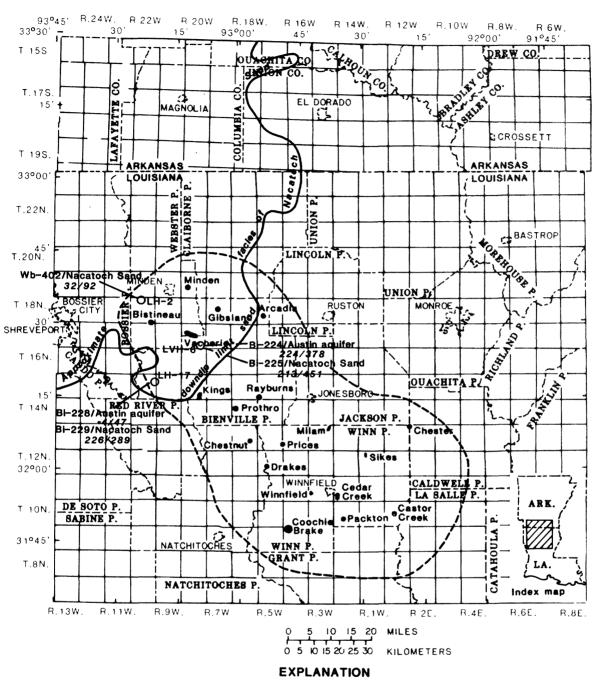
Not water bearing.	Not water bearing.	Saltwater bearing in sand facies.	Not water bearing.	Not water bearing.	Not water bearing.	Not water bearing.	Saltwater bearing.	Not water bearing.	Saltwater bearing.
		Macatoch Sand aquifer				ışışı	Austin aqu		
Marine clay with thin calcareous basal unit. Thickness about 600 ft.	Dark-gray marl, partly chalky. Thickness about 200 ft (60 m).	Shale, sandy, calcareous; sand facies in upper part. Total thickness 250 to 350 ft (75 to 105 m). Sand thickness 100 to 200 ft (30 to 60 m).	White fossiliferous chalk; some marl. Thickness about 50 ft (15 m).	Marl, fossiliferous, chalky. Some glauconitic sand. Thickness 150 to 200 ft (45 to 60 m).	Fossiliferous chalk; marl in lower part. Thickness 200 to 250 ft (60 to 75 m).	Calcareous clay and some sand. Thickness about 100 ft (30 m),	Poorly sorted sands with silt, ash, shale; lignitic, calcare- ous. Thickness about 300 ft (90 m).	Micaceous shale with some micaceous and calcareous sandstone. Thickness about 200 ft (60 m),	Interbedded shales and poorly sorted sands; trace of gravel at base. Thickness less than 200 ft (60 m).
Porters Creek Clay, Clayton Fm.	Arka- delphia Marl	Nacatoch Sand	Sara- toga Chalk	Marl- Marl- Marl-	Аппопа Сћајк	Browns- Marl Marl	Tokio Formation	Spæle Eord Es8le	Woodbine Formation
Midway									
Paleo-				suc	Cretacec	Upper		-	
					taceous	Gre			
					sozoic	∍M ———			

Table 2.--Results of aquifer tests, fiscal year 1980 test-drilling program

Wel	Well No.	ļ	Location	uo	Screened interval (ft	Aquifer	Sand thick-	Date	Yield	Hydraulic
USGS	LETCO	Sec.	(N.)	R. (W.)	below land- surface datum)		ness (ft)	of test (1980)	(gal/min)	conductivity (ft/d)
Wb-401 Wb-402 Wb-403	LH- 2WS LH- 2A LH- 2B	m m m	18 18 18	01 01 01 01	62- 134 1,663-1,795 659- 740	Sparta/Terrace Nacatoch Sand Wilcox-Carrizo	131 120 76	3/17- 3/19 3/24- 3/26 3/13- 3/16	41 29 86	$\frac{1}{3}$
						SITE LH-7				
Bi-217	LH- 7WS	∞	17	2	[477- 549] [579- 590]	Sp art a	89	5/28- 5/30	70	99
Bi-218	LH~ 7A	2	17	5	$\begin{bmatrix} 1,480-1,500 \\ 1,520-1,561 \end{bmatrix}$	Wilcox-Carrizo	55	5/11- 5/13	16	.1
Bi-219	LH- 7B	∞	17	5	$\begin{bmatrix} 943 - & 959 \\ 976 - 1,047 \end{bmatrix}$	Wilcox-Carrizo	91	5/16- 5/18	78	80
						SITE LRH-13				
Bi-220	LRH-13WS	1	17	9	78- 192	Sparta	183	6/24- 6/26	07	1/
Bi-221	LRH-13A	7	14	9	1,047-1,108	Wilcox-Carrizo	133	7/ 9- 7/15	36	7.
Bi-222	LRH-13B	1	14	9	564- 662	Wilcox-Carrizo	76	6/19- 7/22	84	&
	;					SITE LVH-6				
Bi-223	LVH~ 6WS	35	17	80	32- 94	Sparta	06	7/29- 7/31	55	1/
Bi-224	LVH- 6A	35	17	œ	2,537-2,577	Austin	85	9/ 3- 9/ 7	35	2/
Bi-225 Bi-226	LVH- 6B LVH- 6C	35 35	17	∞ ∞	1,831-1,918 909- 970	Nacatoch Sand Wilcox-Carrizo	81 65	8/21- 8/25 7/23- 7/27	83	<.1 7
						SITE LH-17				
Bi-227 Bi-228	LH-17WS LH-17A	12	15	10	294- 428 1,993-2,022	Wilcox-Carrizo Austin	152 26	9/ 9- 9/16 10/20-10/22	6 6	42/
Bi-229	LH-17B	12	15	<u>ස</u>	1,312-1,373	Nacatoch Sand		10/ 7-10/ 9		<.1

1/4quifer properties determined from the tests are much lower than values determined from other tests in the area. Because of water-table conditions, the wells probably were not pumped long enough for the determination of aquifer properties.

2/Because of the large amount of gas present, current methods of aquifer-test analysis probably are not capable of calculating reliable aquifer properties. Permeabilities determined from sidewall cores ranged from 6-1,350 millidarcies.



O LH-2

Approximate boundary of northern Louisiana salt-dome basin

Rayburns
Salt dome and name

Area characterization test-drilling site and number

BI-225/Nacatoch Sand Well number/stratigraphic unit or aquifer 213/451

Altitude of measured head/ measured head corrected to equivalent fresh water head. National Geodetic Vertical Datum of 1929, in feet. (Measured head could be in error because of gas content of water and density differences in water column in well)

Figure 3.--Potentiometric data for the Austin and Nacatoch Sand aquifers.

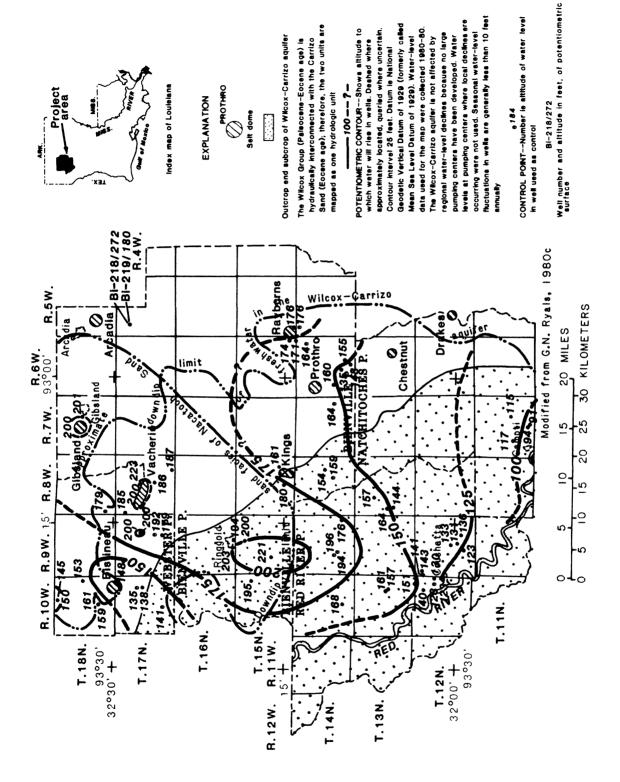
Austin and Nacatoch Sand Aquifers

Five test wells were completed in sands of the Austin and Nacatoch Sand aquifers during the 1980 test drilling; these are the only known hydrologic-test wells completed in the Cretaceous in northern Louisiana. Aquifer testing (table 2) showed that the two units have relatively poor water-transmitting capabilities. Regional potentiometric-surface maps to define gradients of the two units cannot be constructed at this time because of insufficient data. The potentiometric data that are presented in figure 3 indicate that injection wells in the Nacatoch Sand aquifer and petroleum production from the Austin aquifer may have modified the natural system in the aquifers, because head no longer increases with depth. The data indicate a potential for downward flow from the Nacatoch Sand aquifer through the confining layers to the Austin aquifer. ever, site LH-17 (fig. 3) is the only regional site where both the Nacatoch and Austin are screened. Measured water levels may be in error because of the gas content of the water and density differences in the water columns of the test wells. In addition, the sands of the Austin may not be interconnected. Discharge from or recharge to a particular sand bed may not be reflected in other sands of the Austin; thus, waterlevel comparisons may not be valid.

Wilcox-Carrizo Aquifer

Potentiometric data from seven test wells in the Wilcox-Carrizo resulting from the 1980 drilling program were incorporated with other water-level data to construct a generalized potentiometric-surface map of the Wilcox-Carrizo aquifer (fig. 4). The map indicates that the potentiometric surface slopes away from a high south of Ringgold, La. In the part of the study area where the unit contains saline water, potentiometric data are available for only a few points. Most of the potentiometric data for the Wilcox-Carrizo are within the area where the unit contains freshwater and overlies the sand facies of the Nacatoch (fig. 4). Comparing the few data shown in figures 3 and 4, the head in the Nacatoch is at a higher elevation than the Wilcox-Carrizo; thus, upward flow from the Nacatoch, through the confining layer, to the Wilcox-Carrizo may be occurring. Where the sand facies of the Nacatoch is absent, data are not available to compare water levels of the Wilcox-Carrizo and the underlying Austin aquifer.

Ground-water gradients in the Wilcox-Carrizo have been changed locally by injection of oil-field brines. Data from two 1980 test wells may illustrate these effects. Well Bi-218 screened in the lower part of the Wilcox at site LH-7 has a much higher water level than well Bi-219 screened in the upper part of the unit at the same site (fig. 4). Well Bi-219 has a water level that is consistent with regional levels. A nearby injection well is screened in the Wilcox-Carrizo in the same interval as Bi-218. Well Bi-219 is probably not responding to the stress of the injection well at this time because interverning clay beds retard the vertical movement of injection water.



Red Figure 4.--Potentiometric surface of the Wilcox-Carrizo aquifer; Bienville, River, northern Natchitoches, and southern Webster Parishes, La.

Sparta Aquifer

More potentiometric data are available for the Sparta aquifer than for any other water-bearing unit in the study area because of the extent and development of the aquifer. Widespread and intensive pumping of the unit has caused development of numerous cones of depression and a regional lowering of water levels. As shown on the 1980 potentiometric-surface map (fig. 5), major cones of depression are centered at El Dorado, Ark., and Monroe, La. The natural pattern of easterly flow of ground water from outcrop areas to discharge areas in the Mississippi Valley has been significantly modified. Water levels have been lowered below the top of the Sparta in much of northern Louisiana and southern Arkansas. Because of the lack of potentiometric data for the Wilcox-Carrizo, regional flow relations between the Sparta and Wilcox-Carrizo are not well documented. However, because of extensive lowering of water levels in the Sparta, upward flow probably is occurring from the Wilcox-Carrizo toward the Sparta in most of the area.

Under natural flow conditions, movement of water was from the overlying Cockfield Formation (where present) downward through the Cook Mountain confining layer to the Sparta, except in the Mississippi Valley where the Cockfield discharges into the overlying alluvium. The Cook Mountain confining layer contains a discontinuous basal sand bed in parts of the area. Because of the lowering of the water levels in the Sparta, water drains from this sand bed of the Cook Mountain into the Sparta.

MODELING APPROACH AND DATA NEEDS

The multilayered ground-water flow model planned by the Geological Survey for this study will simulate flow in and between aquifers. Regional flow paths and time of travel along any flow path then can be determined under current and projected conditions. Should a contaminant excape from a salt dome, the regional flow model could be used to identify endangered areas and provide information to help cope with the contaminant. The planned model will also provide the basis for a subsequent solute-transport model should one be necessary to solve complex solute-transport problems. Thus, after the regional model is completed, solute-transport modeling techniques, which take into consideration sorptive properties and fluid density, could be incorporated.

The Wilcox-Carrizo, Sparta, and Cockfield aquifers are the principal units to be simulated by the flow model. Because the tentative boundaries of the flow model extend well beyond the salt-dome basin, the Cockfield is included in the model to aid in simulating flow into or out of the top of the Sparta. The Nacatoch Sand may also be an important aquifer to include in the model to aid in simulating flow into or out of the base of the Wilcox-Carrizo. The Austin aquifer is not included in potential modeling plans at this time. Water in the unit, based on data from the 1980 test-drilling program, had a very high gas content and Law Engineering Testing Company (1980c) indicates that there is oil and gas production from the unit. Because the unit may be a multiphase system in the study area, current modeling techniques may not be suitable for simulating flow in the unit.

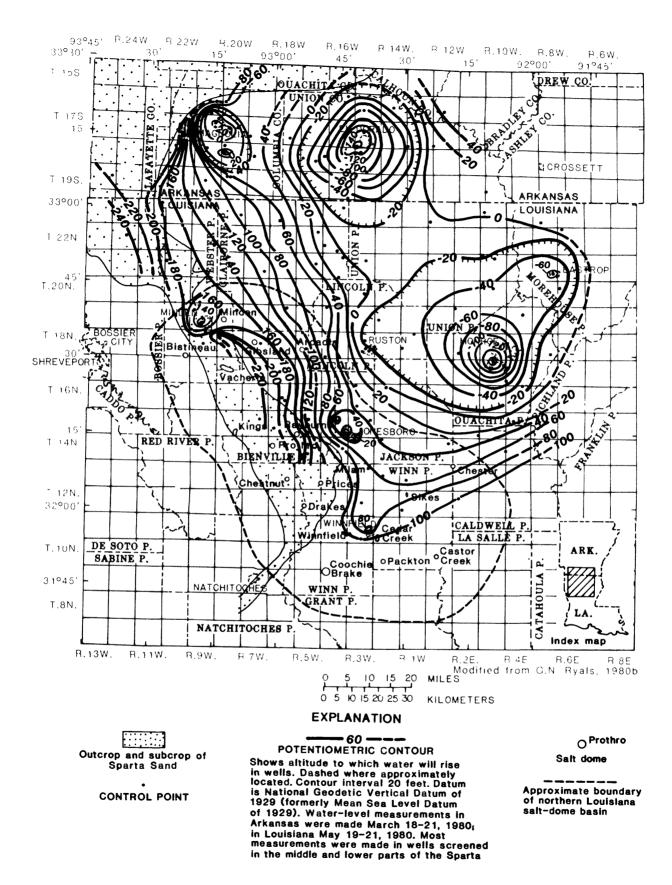


Figure 5.--Potentiometric surface of the Sparta aquifer, northern Louisiana and southern Arkansas, spring 1980.

A three-dimensional model of the Sparta aquifer is being developed to provide a foundation from which to build a detailed model as additional data are collected from regional test holes that are proposed. Additional layers will be added and input into the Sparta model will be refined as more data become available.

Four types of data needed for input to the ground-water flow model and to further develop concepts of the flow system are summarized below. The first is hydraulic properties of the aquifers, such as hydraulic conductivity, transmissivity, and storage coefficient. The second is potentiometric data to define the direction and rate of ground-water flow, both within and between aquifers. The third type of data, hydraulic conductivity of the confining layers, is needed to establish rates of movement between aquifers. No hydraulic conductivity values are known to have been determined for any confining unit in the area. The fourth data category where data is deficient is water chemistry. Ultimately, water chemistry will be very important if a dome is chosen as a repository and solute-transport techniques are used to simulate the ground-water flow system.

Geologic data are available for all of the stratigraphic units, but hydrologic data are sparse. As shown in other sections of the report, hydrologic data are available for areas where the Wilcox-Carrizo, Sparta, and Cockfield aquifers contain freshwater (pl. 2). Electrical logs of wells available in northern Louisiana and southern Arkansas made it possible to define the area of occurrence of freshwater. Hydraulic characteristics of the freshwater section have been determined from test wells and from tests of public-supply and private wells. Most of the pertinent reports that discuss the hydraulic characteristics of the freshwater section are listed in the selected references. Few hydrologic data are available for the saline parts of the Wilcox-Carrizo, Sparta Sand, and Cockfield aquifers and for the Austin and Nacatoch Sand aquifers in the project area.

In 1977 the U.S. Geological Survey in Louisiana proposed a regional test-drilling program to collect the additional data needed for the The program was revised in 1978 in coordination with other proposed drilling. The proposed drilling locations are shown in figure 2; table 3 gives anticipated aquifer depths at the test-drilling sites. Because of the 1980 drilling program and other test wells completed in the area since 1978, sites 4, 11, and 12 can be eliminated from the proposed drilling program. Of the remaining locations, sites 2, 7, 8, 9, and 13 have the highest priority as they are needed to provide data for modeling the Wilcox-Carrizo, Sparta, and Cockfield aquifers. prove to be important in establishing potentiometric gradients in the Nacatoch Sand aquifer from outcrop areas in Arkansas to the northwestern part of the salt-dome basin. Should the Austin aquifer be included in the model, the need for additional information at existing drilling sites and at proposed sites will need to be evaluated.

Table 3.--Depth intervals anticipated for aquifers at proposed regional test-drilling sites

Site	Parish	Location	Aquifer	Aquifer depth in- terval (ft below land-surface datum)
1	Webster	T. 23 N., R. 10 W.	Sparta Sand Wilcox-Carrizo- Nacatoch Sand	700-1,200
2	Union	T. 22 N., R. 2 W.	Sparta Sand Wilcox-Carrizo-	
3	Union	T. 21 N., R. 3 W.	Sparta Sand Wilcox-Carrizo-	
4	Bossier	T. 19 N., R. 12 W.	Wilcox-Carrizo- Nacatoch Sand-	
5	Claiborne	T. 19 N., R. 6 W.1/	Sparta Sand Wilcox-Carrizo- Nacatoch Sand	900-1,600
6	Bienville	T. 16 N., R. 8 W.	Sparta Sand Wilcox-Carrizo- Nacatoch Sand	500-1,100
7	Jackson	T. 17 N., R. 2 W.	Sparta Sand Wilcox-Carrizo	•
8	Ouachita	T. 17 N., R. 3 E.	Sparta Sand Wilcox-Carrizo-	•
9	Richland	T. 17 N., R. 8 E.	Sparta Sand Wilcox-Carrizo-	•
10	Jackson	T. 14 N., R. 4 W.	Sparta Sand Wilcox-Carrizo	
11	Natchitoches-	T. 10 N., R. 10 W.	Wilcox-Carrizo-	- 200- 600
12	Winn	T. 10 N., R. 3 W.	Cockfield Sparta Sand Wilcox-Carrizo-	400-1,100
13	Winn	T. 12 N., R. 2 W. <u>2</u> /	Cockfield Sparta Sand Wilcox-Carrizo	400-1,100
14	Franklin	T. 12 N., R. 6 E.	Cockfield Sparta Sand Wilcox-Carrizo-	1,200-2,000

^{1/}NW/4 of T. 19 N., R. 6 W. or SE/4 of T. 19 N., R. 7 W. 2/Changed from T. 12 N., R. 1 E. in 1978 U.S. Geological Survey drilling proposal.

SUMMARY AND CONCLUSIONS

The method of approach that is being used to select a salt dome for a repository, studying smaller and smaller areas in greater detail, has emphasized the collection of data near potential repository sites. Progress on developing a ground-water flow model to simulate the regional flow system under current and projected conditions has been slowed by a lack of data, particularly for the aquifers that contain saline water. Planned regional geohydrologic studies will define the direction and rate of movement of ground water, which would be necessary information should a contaminant escape from a repository in a salt dome.

In the northern Louisiana salt-dome basin, the interval between land surface and the potential maximum repository depth of 3,000 ft includes the Tokio Formation and Brownstown Marl (Austin aquifer in this report), Nacatoch Sand, Wilcox Group, Carrizo Sand, Sparta Sand, and Cockfield Formation, all of which contain important regional aquifers. The Wilcox Group and Carrizo Sand are interconnected and are treated as one hydrologic unit. Freshwater occurs in parts of the Wilcox-Carrizo, Sparta, and Cockfield aquifers, but the Austin and Nacatoch Sand aquifers contain saline water throughout the basin. Clay and marl units confine the aquifer and produce artesian conditions, except in the outcrop where water-table conditions prevail.

Conceptually, under natural conditions regional ground-water flow is from recharge areas of the aquifers downdip toward the southeast. Natural flow patterns have been altered because of man's activities since the early 1900's. Rates and directions of movement have been changed because of freshwater-production wells in the Wilcox-Carrizo, Sparta, and Cockfield aquifers and because of injection wells in the saline aquifers and the saline part of the Wilcox-Carrizo aquifer. Except for the Sparta, potentiometric data are not available to define regional directions of water movement in the aquifers.

Data from test-drilling in 1980 provided some indication of the potential for vertical movement. Water may be moving downward from the Nacatoch Sand aquifer to the Austin aquifer and upward from the Nacatoch (where present) to the Wilcox-Carrizo. Because of extensive development of the Sparta aquifer, water levels have been lowered below the top of the unit in many areas; thus, water probably moves upward from the Wilcox-Carrizo in most of the area. Under natural conditions water movement probably was from the Cockfield to the Sparta, except in discharge areas in the Mississippi River valley.

Concepts of present conditions of regional flow are based on sparse data. Data are available for the freshwater parts of the aquifers; however, for the saline parts, the Austin and the Nacatoch Sand aquifers, few data are available. Data needed to refine the conceptual model and to develop the ground-water flow model include aquifer properties, potentiometric data, hydraulic conductivity of confining beds, and water chemistry. A comprehensive regional test-drilling program has been proposed to obtain the needed data. Sufficient data are available for

the Sparta aquifer to develop a limited ground-water flow model. As data from the regional test-drilling program become available, the Sparta model can be expanded and refined to define and evaluate the regional flow system.

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Table 4.--Chemical analyses of water from fiscal year 1980 test wells

Well No.	Lo Sec.	catio T.	n R.	Site No.			Date of sample (1980)	Specific conductance (umhos)	pH (units)	Temperature (°C)	<pre>Color (platinum-cobalt units)</pre>	Hardness (mg/L as CaCO3)	Hardness, noncarbonate (mg/L CaCO ₃)	Calcium, dissolved (mg/L as Ca)	Magnesium, dissolved (mg/L as Mg)	Sodium, dissolved (mg/L as Na)
								SPARTA	A QUI	FER						
WB-4011/	3	18N	10W	LH-			3-18	320	6.2	19.5	5	68	0	17	6.2	36
BI-217 BI-220	8 1	17N 14N	5W 6W	LH- LRH-1			5-29 6-25	250 76	6.2 5.1	19.0	30 5	54 22	0	17 6.3	2.8 1.5	25 3.
BI-223	35	17N	8W	LVH-			7-30	20	4.6	18.0	50	4	Ö	.9	.3	4.
							WI	LCOX -CAF	RIZO	AOUIFER	 !					
WB-403	3	18N	10W	LH-	2	740	3-15	6,500	6.7	27.0	5	120	0	33	10	1,300
BI-218	5	17N	5W	LH-			5- 7	12,900	7.9	31.5	0	89	0	20	9.3	2,700
BI-219 BI-221	8 1	17N 14N	5W 6W	LH- LRH-1			5-17 7-10	4,460 719	8.0 8.3	22.0 29.5	30 30	21 19	0	5.0 6.6	2.0 .5	960 150
BI-222	î	14N	6W	LRH-1			6-20	656	8.3	25.0	60	8	0	2.7	.4	160
BI-226	35	17N 15N	8W	LVH-	6 9	70	7-24	7,700	8.4	28.5	5	42	0	11	3.5	1,500
BI-227	12	13N	10W	LH-1	. / 4	28	9-10	340 ACATOCH	7.6	22.5	0	98		30	5.6	33
WB-402	3	18N	10W	LH-	2 1,7	205	3-23		7.2	38.5	20	2,700	2,500	650	260	17,000
BI-225	35	17N	8w	LVH-	6 1,9	18	8-22	183,000	7.0		40	4,300	4,100	1,100	370	85,000
BI-229	12	15N	10W	LH-	1,3	373 1	0- 7	45,400	7.9	28.5		1,800	1,600	440	160	10,000
								AUSTIN	A QUI	FER						
BI-224 BI-228	35 12	17N 15N	8W 10W	LVH-		597 022 1	9- 4 0-21	154,000 66,200	6.6 7.5	42.0 33.0	15	6,600 2,700	6,500 2,600	1,800 800	510 180	39,000 15,000
Well No.	Potassium, dissolved (mg/L as K)	Alkalinity, field	(mg/L as CaCO ₃)	Bicarbonate (mg/L as HCO_3)	Carbonate (mg/L as ${\rm CO}_3$)	Carbon dioxide, dissolved (mg/L as ${ m CO}_2)$	Sulfate, dissolved mg/L as SO.)	Chloride, dissolved		Fluoride, dissolved (mg/L as F)	Silica, dissolyed (mg/L as SIO ²)	Solids, residue at	180°C, dissolved (mg/L)	Solids, sum of constituents, dissolved (mg/L)	Iron, dissolved (µg/L as Fe)	Manganese, dissolved (µg/L as Mn)
							SPA	RTA AQUI	FER	Cont inu	ed					
WB-4011/	2.7			130	0	131		.1	40	0.1	37		196	206	760	30
BI-217 BI-220	3.9 2.8		2	100 40	0	91 508	35	.0	7.9		60 26		193 68.0	199 64.0	2,500 1,200	70 20
BI -223	1.1		8.0	10	Ö	49		.1	3.2		16		37.0	33.0	40	2
						V	ILCOX-	CARRIZO	A QUIF	ERCon	tinued					
WB-403	8.8			438	0	140			000	.4	15	3	,360	3,590	90	80
BI-218	12	51		626	0	13	28		000	1.6	15			7,100	190	40
BI-219 BI-221	4.4 1.9			1,210 310	0	19 2.5	24	.1	880 61	3.0 .2	13 17	2	,460 441	2,460 412	150 50	0 10
BI-222	1.3			408	0	3.1		.3	10	.3	15		403	388	1,600	90
BI-226	5.5	40	4	493	1	3.0		.6 2,	100	.8	60	3	,760	3,930	150	30
BI-227	2.1	16	5					.6	8.0		28	 	215	206	260	70
							NACATOO	H SAND A	QUIFE	RCont	inued					
							7.7	30	000	.5	14	51	,100	48,200	2,500	100
WB-402	77	15		192	0	19										
BI-225	31	15 14 12	6	192 178	0 - -	19 28 3.0	2,400	130,	000	.2	4.	1 223	,000	219,000	18,000	450
		14	6	178	-	28	2,400	130	000	.7	11	1 223				
BI-225	31	14	6	178	-	28	2,400	130, 17, rin Aqui	000	.7	11	1 223 28	,000	219,000	18,000	450 170

 $\underline{1}/\text{Terrace deposits also screened in this well.}$

Table 5.--Trace-metal analyses of water from fiscal year 1980 test wells

Well No.	Lo Sec.	catio T.	n R.	Site No.	Depth of well (feet)	Date of Sample (1980)	Aluminum, dissolved (µS/L as Al)	Antimony, dissolved (bg/L as Sb)	Arsenic, dissolved (µg/L as As)	Barium, dissolved (µg/L as Ba)	Beryllium, dissolved (ug/L as Be)	Cadmium, dissolved (µg/L as Cd)	Chromium, dissolved (ug/L as Cr)
						:	SPARTA A	QUIFER					
WB-4011/ BI-217 BI-220 BI-223	3 8 1 35	18N 17N 14N 17N	10W 5W 6W 8W	LH- 2 LH- 7 LRH-13 LVH- 6	134 590 192 94	3-18 5-29 6-25 7-30	10 0 0 0	1 0 0 0	0 1 1 0	100 90 100 20	0 <1 <1 <1	0 <1 <1 <1	0 0 10 10
						WILC	OX -CARRI	ZO AQUIFE	R 				
WB-403 BI-218 BI-219 BI-221 BI-222 BI-226 BI-227	3 5 8 1 1 35 12	18N 17N 17N 14N 14N 17N 15N	10W 5W 5W 6W 6W 8W 10W	LH- 2 LH- 7 LH- 7 LRH-13 LRH-13 LVH- 6 LH-17	740 1,561 1,047 1,215 662 970 428	3-15 5- 7 5-17 7-10 6-20 7-24 9-10	10 30 10 0 10 0 10	2 2 0 0 0 0 1	0 0 0 1 1 0 0	700 800 100 80 40 100 200	0 10 0 1 <1 10 <1	0 0 0 <1 <1 1 1	0 10 10 10 10 10
WB-402 BI-225	3 35	18N 17N	10W 8W	LH- 2 LVH- 6	1,795 1,918	3-23 8-22	10 0	3 1	0 0	8,800 1,200	20	0 0	30 190
BI-229	12	15N	10W	LH-17	1,373	10- 7	o	1	1	7,000	10	0	40
							AUSTIN A	QUIFER					
BI-224 BI-228	35 12	17N 15N	8W 10W	LVH- 6 LH-17	2,697 2,022	9- 4 10-21	0	0 1	0	7,500 6,00 0	0 10	0	60 70
Well No.		Cobalt, dissolved (µg/L as CO)	Copper, dissolved (µg/L as Cu)	Iron, dissolved	(μg/L as Fe)	Lead, dissolved (µg/L as Pb)	Lithium, dissolved (µg/L as Li)	Manganese, dissolved (µg/L as Mn)	Mercury, dissolved (µg/L as Hg)	Molybdenum, dissolved (µg/L as Mo)	Nickel, dissolved (µg/L as Ni)	Silver, dissolved (µg/L as Ag)	Zinc, dissolved (µg/L as Zn)
						SPARTA	AQUIFER	RCont inu	ied				
WB-401 <u>1</u> / BI-217 BI-220 BI-223		0 <3 <3 <3	0 0 1 0	2,5 1,2		0 0 3 1	20 20 10 6	30 70 20 2	0.0 .0 .1	0 <10 <10 <10	2 0 7 0	0 1 0 0	130 440 610 490
					WI	LCOX -CAR	RIZO AQU	JIFERCor	tinued				
WB-403 BI-218 BI-219 BI-221 BI-222 BI-226 BI-227		0 1 0 <3 <3 0 <3	0 0 0 0 0	1 1 1,6	90 90 50 50 00 50	0 0 2 1 0 2	70 200 50 10 8 70 20	80 40 0 10 90 30 70	.2 .3 .2 .0 .1 .2	0 1 1 <10 <10 1	2 3 2 0 4 0	0 0 0 0 0	80 240 40 <3 <3 40 80
					N.A	CATOCH	SAND AQU	IFERCon	tinued				
WB-402 BI-225 BI-229		0 0 3	1 0 5	2,5 18,0 4,3	00	0 16 2	820 240 640	100 450 170	.0 .1 .4	0 1 2	1 2 2	0 0 0	2,400 8,400 560
BI-224 BI-228		0 2	0	7,5 8,1		1 1	700 320	1,000 370	.5 1.0	0	0	0	1,800 1,600

 $\underline{1}/\text{Terrace deposits also screened in this well.}$

Table 6.--Concentrations of radioactive elements in water from fiscal year 1980 test wells

Well No.	Lo Sec.	catio T.	n R.	Site No.	Depth of well (feet)	Date of sample (1980)	Gross alpha, dissolved (pCi/L as U-natural)	Gross alpha, suspended total (pCi/L as U-natural)	Gross beta, dissolved (pCi/L as Cs-137)	Gross beta, suspended total (pCi/L as Cs-137)	Gross beta, dissolved (pC1/L as Sr/Yt-90)	Gross beta, suspended total (pCi/L as Sr/Yt-90)
							AQUIFER					
WB-401 <u>1</u> /	3	18N	10W	LH- 2	134	3-18	2.3		4.3		4.4	
BI-217	8	17N	5W	LH- 7	590	5-29	<1.8		2.0		1.9	
BI-220 BI-223	1 35	14 N 17 N	6W 8W	LRH-13 LVH- 6	192 94	6-25 7-30	<.4 <.3		3.0 1.0		2.8 .9	
								TEED				
							RIZO AQU	17EK				
WB-403 BI-218	3 5	18N 17N	10W 5W	LH- 2 LH- 7	740 1,561	3-15 5- 7	<49 <82		<39 <80		<40 <77	
BI-219	8	17N	5W	LH- 7	1,047	5-17	<26		<20		<19	
BI-221	1	14N	6W	LRH-13	1,215	7-10	<4.1		<3.5		<3.4	
BI-222	1	14N	6W	LRH-13	662	6-20	<3.1		3.3		3.1	
BI-226 BI-227	35 12	17N 15N	8W 10W	LVH- 6 LH-17	970 428	7-24 9-10	<66 <2.7		<46 <1.6		<44 <1.6	
					NA C	ATOCH :	SAND AQU	IFER				
₩B-402	3	18N	10W	LH- 2	1,795	3-23	< 750		<560		<580	
BI-225	35	17N	8W	LVH- 6	1,918	8-22<	<2,400	•	<2,200		<2,100	
BI -229	12	15N	10W	LH-17	1,373	10- 7	310	6.7	<300	1.4	<280	1.4
						AUSTIN	AQUIFER					
BI-224	35	17N	8w	LVH- 6	2,697		<1,400		<1,200		<1,100	
BI -228	12	15N	10W	LH-17	2,022	10-21	610	19 	<520	8.0	< 500	7.6
Well No.	Castum 137 dis-	solved (pCi/L)	Potassium 40, dis-	solved (pCi/L as K40) Potassium 40,	total (pCi/L) Radium 226, dis-	solved, planchet count (pCi/L)	Radium 226, dissolved, radon method (pCi/L)	Strontium 90, dissolved (pCi/L)	Tritium, total (pG1/L)		Uranium natural, dissolved (µg/L as U)	<pre>Uranium, dissolved, extraction (µg/L)</pre>
					SPART	A AQUI	FERCont	inued				
WB-401 <u>1</u> /	< 1	.0					0.43	<0.4	<24		<0.6	
BI-217		.0					.16	< .4	<20		<.6	
BI-220 BI-223		.0					.15 .11	< .4 < .4	<30 37		<.6 	<.70
					LCOX -CA	RRIZO	A QUIFER-	-Cont in	ıed			
WB-403	< 2	.0	_				.83	< . 4	<24		<.6	
BI-218		.0					2.6	< .4	<20		<.6	
BI-219		.0		3.	3 -		.55	< .4	<20			<.50
BI-221	< 1	.0					.12	<.4	<30		<.6	
BI-222		.0					.04	< .4	<30		<.6	
BI -226		.0					.79	< .4	<30			<.60
BI-227	< 1	.0				- 	.30		<10			<.60
				N.	ACATOOH	SAND A	QUIFER	Cont inu	ed			
WB-402	< 2	.0					67	1.5	<24			<14
BI-225 BI-229		.0		 28			120 26	.7	<10 <20		<1.6	<5.0
D1-77A							-	<.4			<1.6	
							FERCon				 	
BI -224		.0			1 	L40 65	65	3.5	<10 <20		<1.6	<5.0
BI -228	· · ·			· · · · · · · · · · · · · · · · · · ·		65		<.4	<20		<1.6	

 $\underline{1}/\text{Terrace}$ deposits also screened in this well.

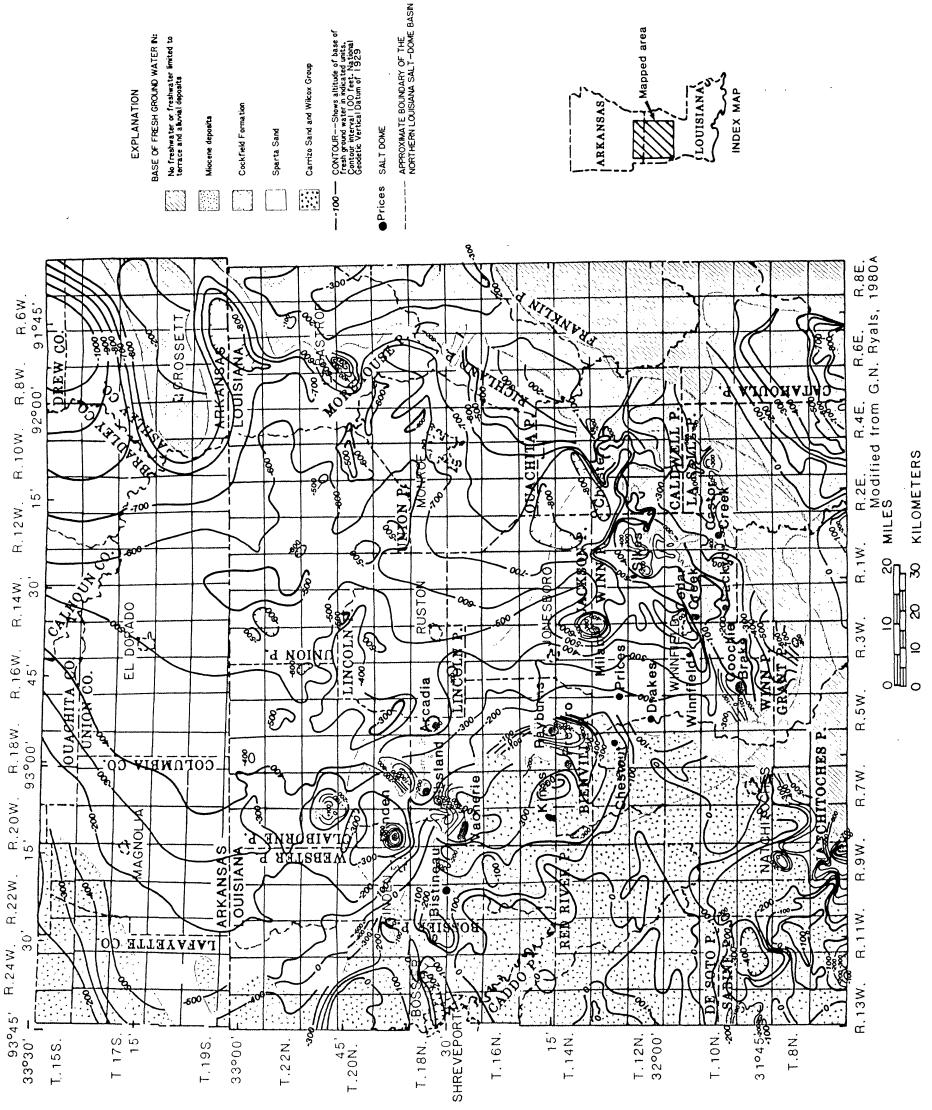


PLATE 2. MAXIMUM DEPTH OF OCCURRENCE OF FRESHWATER, NORTHERN LOUISIANA SALT-DOME BASIN AND VICINITY.

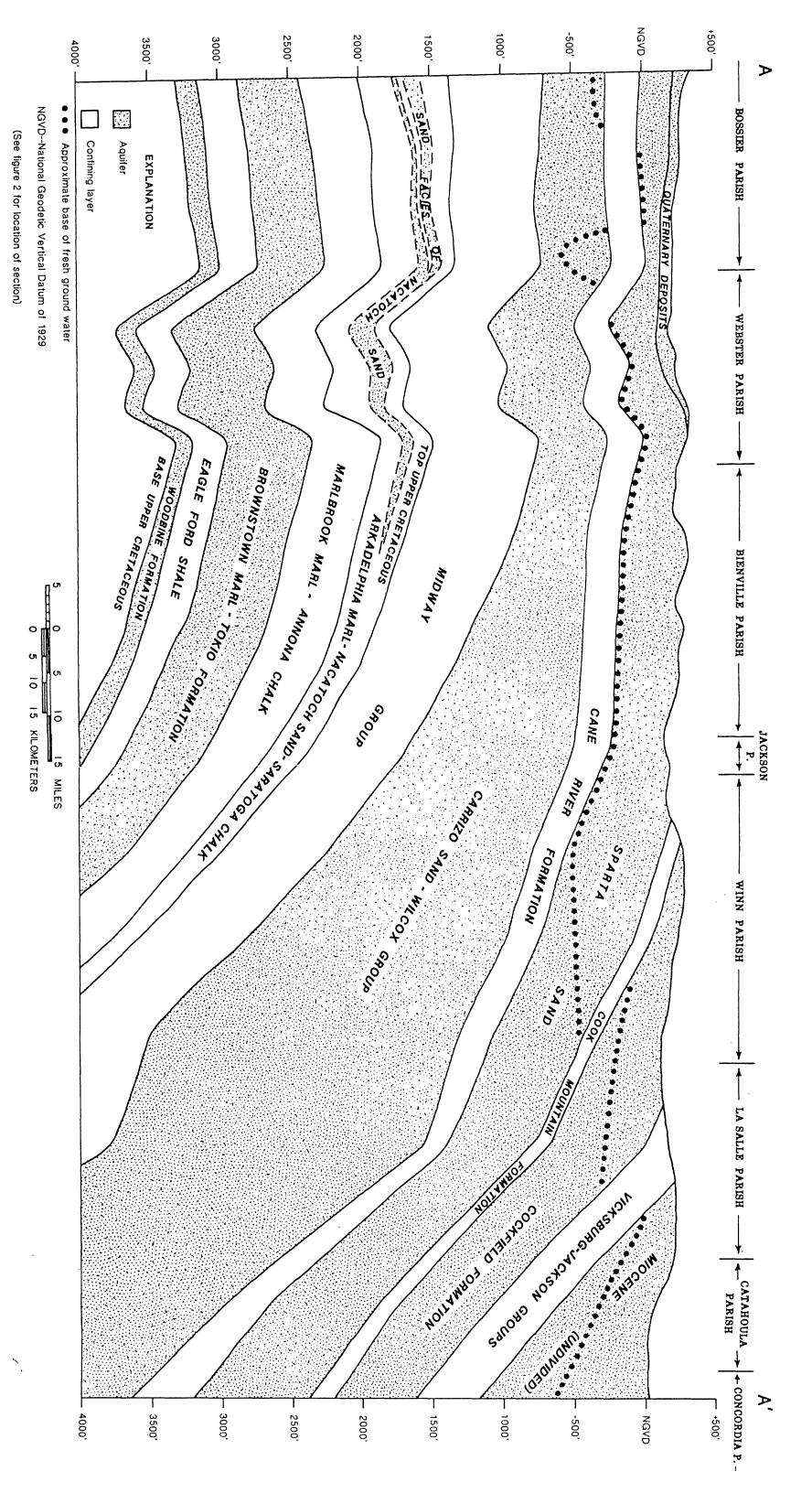


PLATE 1. GENERALIZED NORTHWEST-SOUTHEAST GEOLOGIC SECTION, NORTHERN LOUISIANA SALT-DOME BASIN.